

# Relationship between the Expansion of Mountain Lakes and Their Catchment Mountains on the Qinghai-Tibet Plateau

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Inland closed lakes are good indicators of changes of regional climate and hydrology [1, 2]. Based on the 1:500,000 topography map (1976 edition), the authors chose 5 typical regions and respectively summarized many characteristics of 12~15 lakes and their basins, and in conjunction with the 1:100,000 topography map (1970 edition) the author also investigated rivers, snow lines, glaciers and the climate characteristics in these basins, and researched the relationship between mountains and lakes in different regions, also tried understanding the difference mechanism of expansion scope of mountain lakes during the warm-wet period.

## Hydrological balance of mountain lakes in a plateau

Generally, water balance equation of closed lake basins is shown as the following:

$$P_B S_B + P_L S_L + G_W = E_L S_L + E_B S_B \quad (1)$$

Among which,  $P_B$  and  $P_L$  respectively denote the basin land (excluding glacier, the same below) and the lake-level average precipitation,  $S_B$  and  $S_L$  respectively denote the basin land area and lake area,  $G_W$  denote the glacial meltwater,  $E_L$  and  $E_B$  respectively denote the lake-level evaporation and basin land-level evaporation.

If the equation is:

$$P_B S_B - E_B S_B + G_W = (E_L - P_L) S_L \quad (2)$$

And  $(E_L - P_L) S_L$  of which can be shown as  $\langle RS \rangle$ ,  $\langle RS \rangle$  denotes basin mountain runoff at different heights:

$$\langle RS \rangle = R_1 S_1 + R_2 S_2 + \dots + R_n S_n \quad (3)$$

Among which,  $R$  denotes runoff depth,  $S$  denotes the mountain area at different heights. So, Lake area and mountain area at different heights have the following relationship:

$$(E_L - P_L) S_L = R_1 S_1 + R_2 S_2 + \dots + R_n S_n \quad (4)$$

Introducing into lake supply coefficient  $K = S_B / S_L$ <sup>[3]</sup>, we can elicit the following linear relationship between  $K$  and the mountain area at different heights:

$$(E_L - P_L) K = R_1 S_1 / S_B + R_2 S_2 / S_B + \dots + R_n S_n / S_B \quad (5)$$

While the reciprocal of lake supply coefficient  $K_V = 1/K = S_L / S_B$  can denotes the static water maintained by unit basin land surface.

Therefore, from the relationship between the lake area and the mountain area at different heights we can analyze the characteristics of hydrological flow & store and lake expansion of closed lake basins.

## Relationship between lakes and the hydrological characteristics of basins of typical regions in the Plateau

Based on temperature and precipitation, we divided the plateau into 5 typical regions (Fig.1): A) NamCo-Zige-tang Co Region, B) Zhari Namco -Angzi Co Region, C) Akesaiqin Lake-Guozha Co Region, D) Maerguochaka-Duogerenco QiangCo Region, E) Kusai Lake-Wulan Wula Region; and summarized many characteristics of 12~15 closed basins with similar climate in various regions (Table 1).

These basins are basically classified into two types: 1) lowland basins type (remarked as L type), in the basins with meek and open valley lowlands, the mountains with the height of more than 1,000 m (above the present lake level,

except special instruction, the same below) take up less than 1‰, such as Zigetang Co, Yong Co and Gouren Co. 2) Mountain basins type (remarked as M type), in this kind of basins, the mountains with the height of more than 1,000 m take up above 1‰. According to that the mountains have slope break or not, and the height of slope break, basins can be subdivided into four types: (1) Ma type, with 500~750m height of slope break, namely, the lowlands with the height of less than 500~750m are meek while the mountains with the height more than 500~1000m are much steeper, such as Zhabu Yechaka Lake, Jiasa Co, Anggu Co, Guozha Co and Peigu Co, etc. The mountain area with the height of 500~1000m in the basins is larger. (2) MB type, with 1000~1500m height of breach, namely, middle-low mountains with the height less than 1000~1500m are meek while the high mountains with the height of more than 1000~1500m are much steeper, such as Dangqiong Co, Muco Bingni Lake, Dangreyong Co, Dongbu Co and Bangda Co, etc. The mountain area with the height of 1000~1500m in the basins is larger. (3) Mc type, there are slop breaks with the above two heights, such as Suona Lake, Taiyang Lake, Longmu Co, Kekexili Lake, etc. And (4) Md type, there is no obvious slope breaks, and some lakeshore plains are narrow such as Bangong Co; some are opener such as Nam Co, Zhari Namco, Akesaiqin Lake, Zeco and Jiezechaka Lake. In other words, some basin mountains have a plateau surface of relative meek gurgitation, such as Ma, Mb, Mc-type basins, while Md-type basin has no obvious plateau surface.

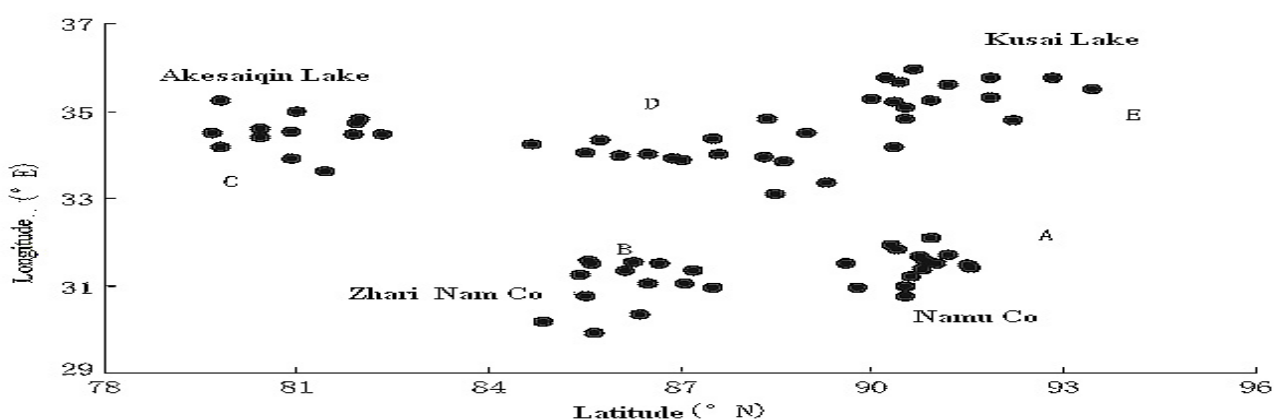
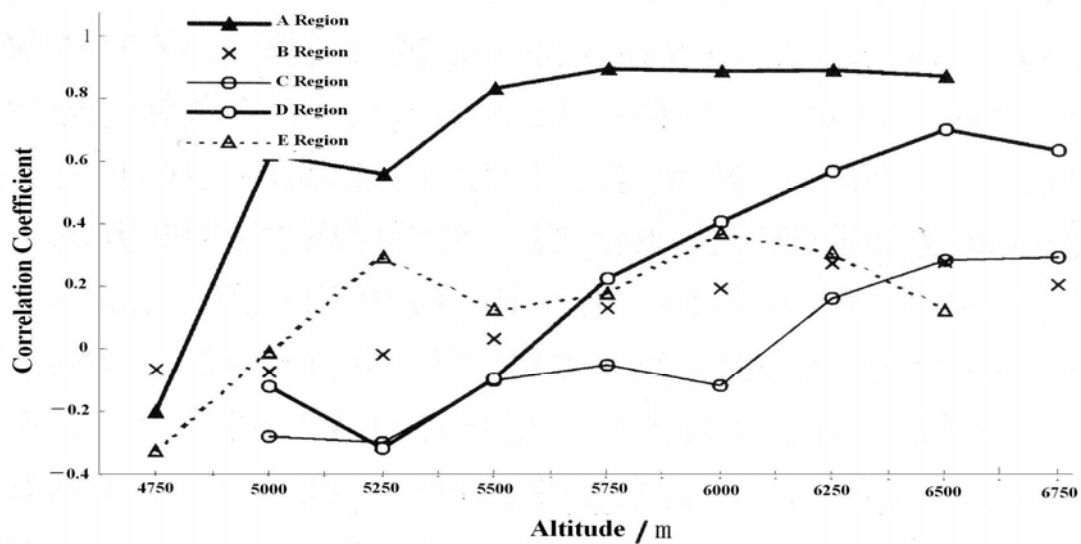


Figure 1, Distribution of Lakes mentioned in this paper on the Qinghai-Tibet Plateau

Table 1 Information of various typical closed lake basins on the Qinghai-Tibet Plateau

| Region | Temperature | Annual Precipitation | Height of Snow Lines | Lake-level Height | Generation Height of Perennial River | Disappeared River (Initial Height) |
|--------|-------------|----------------------|----------------------|-------------------|--------------------------------------|------------------------------------|
| A      | 0~-2?       | 400~450mm            | 5500~5750m           | 4500~4750m        | 5200~5500m                           | N/A                                |
| B      | 0~-2?       | 140~200mm            | 5800~6000m           | 4500~5200m        | 5200~5700m                           | N/A                                |
| C      | <-4?        | 140~200mm            | 5800~6000m           | 4800~5100m        | 5700~6000m                           | 5300~4800m                         |
| D      | <-4?        | <75mm                | 5800~6000m           | 5000~5200m        | 5800~6000m                           | 5300~5000m                         |
| E      | <-2?        | 120~160mm            | 5400~5500m           | 4800~5200m        | 5200~5500m                           | 5000~4800m                         |

These basins are basically classified into two types: 1) lowland basins type (remarked as L type), in the basins with meek and open valley lowlands, the mountains with the height of more than 1,000 m (above the present lake level, except special instruction, the same below) take up less than 1‰, such as Zigetang Co, Yong Co and Gouren Co. 2) Mountain basins type (remarked as M type), in this kind of basins, the mountains with the height of more than 1,000 m take up above 1‰. According to that the mountains have slope break or not, and the height of slope break, basins can be subdivided into four types: (1) Ma type, with 500~750m height of slope break; (2) MB type, with 1000~1500m height of breach; (3) Mc type, there are slop breaks with the above two heights; And (4) Md type. Md-type basin has no obvious plateau surface. while some basin mountains have a plateau surface of relative meek gurgitation, such as Ma, Mb, Mc-type basins.



**Figure 2 Correlation coefficients of  $1/k$  and mountain area at various heights in various regions at the 0.01 confidence level**

The correlation (shown as Drawing 2) between  $K_v$  and the mountain area at different heights in various basins is approximately the same. With the heightening of mountains, related coefficient is bigger and bigger, but the related coefficient of the mountain with the height of more than 6,250m and perennially covered by ice and snow is smaller.

#### **Change of mountain lakes of the plateau during the global warm-wet period since 40 ka B.P**

According to marine (deep-sea) oxygen isotope records, since 40 ka B.P, the world has undergone three typical climate stages: the interstadial of the last glacial---the last glacial maximum stage---Holocene (interglacial stage)<sup>[4]</sup>. Through the comprehensive analysis of existing information<sup>[5-21]</sup>, the closed lakes during the warm-wet period in the interglacial stage of Qinghai-Tibet Plateau have the following expansion characteristics (shown as Table 2 and 3, ellipsis).

At interglacial stage of 40~23 ka B.P, mountain lakes significantly expanded, and at the glacial maximum stage, paleolake area was several times even dozens of times<sup>[21, 22]</sup> of present area or the paleolake became exorheic lake<sup>[20]</sup>.

At Holocene (8.0~5.0 ka B.P), lakes in the plateau entered into another expansion period, but totally, the expansion range was far less than that at 40~23 ka B.P.

#### **Discussions**

Based on most of lakes in Qinghai-Tibet Plateau since 40 ka B.P and Hydrological Effects of lake expansion under the orbital-cycle global-warming period (especially at 40~23 ka B.P), scholars have deep researches<sup>[14,19-22]</sup> on the expansion scope, change and mechanism of regional climate environment, etc.

However, based on basin type, lake status quo and expansion scope, precipitation increase of the plateau in warm-wet period was more likely to happen in the middle and high mountains, instead of hill and lowland. Especially, middle and the high mountains (with the height>500~750m) have larger area, and their lake expansion scope is greater. So, L-type basin has biggest modern lake supply coefficient (in comparison with others in the same region, the same below), while the lake expansion scope is very little. For basins of the highest mountain (with the height>1500~2000m), middle and high mountains with the height>500~750m have greater area (absolute area and the proportion), smaller modern lake supply coefficient and very obvious lake expansion.

The above situation shows that the precipitation of middle and high mountains is more than hill and lowland, which also may be caused by the following three factors:1) The higher the mountain is, the lower its temperature is, and the fainter the evaporation is; after flowing to and storing in a lake, the cold runoff will cool the lake in certain degree, and also reduce the lake evaporation, which makes the lake shrinkage caused by lake evaporation slowly.2) The mountain's runoff depth value is bigger than hill's and lowland's under the same other conditions, which makes hydrographic flow and

store of lake increased. And 3) When the climate changes to the warm-wet period making the regional precipitation increased, the lake expansion scope of middle and high mountains' basins is more obvious than hill's and lowland's. So, when rebuilding paleo-hydrography through exploiting paleolakes site, the regional wet degree of ancient climatic environment will be amplified artificially to some extent.

## Conclusions

1) There is close relation between the mountain lakes area on the plateau and the basin mountain (the height in the east of the plateau >25m, and in the central and the west of the plateau >500~750m). The mountain's precipitation is the main source of lake-water, and the quality (low temperature) and quantity are the main factors for sustaining the existence of a lake which keeps a certain scope steadily.

2) In Qinghai-Tibet Plateau, the lake expansion scope at middle and high mountain basins is more obvious than hill's and lowland's in global warm-wet period and under the climate changing from cold to warm period.

3) At 40~23 ka B.P, the mountain lakes on the plateau had unusually expansion, which may be caused by a mass of low-temperature water formed by increased monsoon precipitation flowed to and stored in the lake during warm period. And the water increase is more obvious at mountain areas.

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## Reference

- [1] Ge Yu, 1997. Studies on global late Quaternary lake-level data bases [J]. *Journal of Lake Science*, 9(3): 193-202 (in Chinese with English abstract).
- [2] Boqiang Qin. 1993. Study on the impact of climate change on the Asian inland land—past present and future [D]. Ph.D. Dissertation of Chinese Academy of Science, 1993: 60-61 (in Chinese with English abstract).
- [3] Fan Yunqi, 1983, the supply coefficient of interior lakes in Xinjiang [J], *Oceanologia Et Limnologia Sinica*, 14(2):118-127 (in Chinese with English abstract).
- [4] Prell, Warren L.; Imbrie, John; Martinson, Douglas G.; et al. 1986. Graphic Correlation of Oxygen Isotope Stratigraphy Application to the Late Quaternary [J]. *Paleoceanography*, vol. 1, issue 2, pp. 137-162.
- [5] Huang Cixuan Zhang Qingsong Liu Futao, 1989. A Preliminary Study of Paleovegetation and Paleoclimate in the Later period of Late Pleistocene in the Bangongcuo Lake Region lake region of Xinjiang [J]. *Journal of Natural Resources*, 4(3):247-253 (in Chinese with English abstract).
- [6] Li Shijie, Zheng Benxing, Jiao Keqin, 1991. Preliminary Research on Lakes Evolution of the Southern Slope of West Kunlun Mountains [J]. *Scientia Geographica Sinica*, 11(4): 306-313(in Chinese with English abstract).
- [7] Hu Dongsheng. 1995. Lakes Evolution in Kekexili region [J]. *Arid Land Geography*, 18(1): 60-67 (in Chinese with English abstract).
- [8] Chen Kezao, Bowler J M, Kelts K, 1990. Climate Change Since 40 kaBP in Qingzang plateau[J]. *Quaternary Sciences*, 10(1):21-31(in Chinese with English abstract).
- [9] Zheng Mianping, Liu Junying, Qi Wen. Climate Change Research Based on Salt Sedimentary Stratigraphy During the Last 40ka in Xizang Plateau [C]// Zheng Mianping. Salt Resource and Environment and Global Change. Beijing: Geological Publishing House, 1996. 6-20(in Chinese with English abstract)
- [10] Zhao Xitao, Zhu Dagang, Yan Fuhua, et al. 2003. Climatic Change and Lake-level Variation of Nam Co, XiZan Since the last Interglacial Stage[J]. *Quaternary Sciences*, 23(1): 41-52(in Chinese with English abstract).
- [11] Jia Yulian, Wang ShuMing, Wu YanHong, et al, 2003. Preliminary Study of Lake Evolution and Precipitation of Zingetang Co and Coe Basins, Central QingHai-XiZang Plateau, Since 24 kaBP [J]. *Oceanologia Et Limnologia Sinica*, 23(5):283-294(in Chinese with English abstract).
- [12] Li Bingyuan. Lake Evolution---the Main Characteristics of Environmental Change in Xiang Plateau [C]// Uplift and Environment Change During Late Cenozoic era in Xizang Plateau (eds. Shi Yafeng, Li Jijun,, Li Bingyuan). Guangzhou: Guangdong Science and Technology Press, 1998, paper 331-372 (in Chinese with English abstract).
- [13] Wang Fu bao, 1987. Climatic changes in the Qinghai-Xizang(Tibetan) region of China during the Holocene [J]. *Quaternary Research*, 28(4): 50-60.
- [14] Zheng Mianping, Xiang Jun. Salt Lake in Xizang Plateau [M]. Beijing: Beijing Science and Technology Press, 1989, paper 120-353 (in Chinese with English abstract).
- [15] F. Gasse, M. Arnold, J. C. Fontes, M. Fort, E. Gibert, A. Huc, Li Bingyuan, Li Yuanfang, Liu Qing, F. MÃ©liÃ¨res, E. Van Campo, Wang Fubao, Zhang Qingsong, 1991. A 13,000-year climate record from western Tibet [J]. *Nature* 353, 742-745
- [16] Comprehensive Science Expedition in Xizang Plateau from Chinese Academy of Science. Quaternary Geology in Xizang [M]. Beijing: Science Press, 1983, 1-251(in Chinese with English abstract).
- [17] Li Shuaike, 1992. Closed Lakes Level Variation and their implications in middle Kunlun Mountain [J]. *Journal of Lake Sciences*, 4(1): 19-30 (in Chinese with English abstract)
- [18] J. Ch. Fontes, F. MÃ©liÃ¨res, E. Gibert, Liu Qing and F. Gasse, 1993. Stable isotope and radiocarbon balances of two Tibetan lakes (Sumxi Co, Longmu Co) from 13,000 BP[J]. *Quaternary Science Reviews*, 12(10): pages 875-887.
- [19] Li Bingyuan, 2000. The last greatest lakes [J] *Acta Geographica Sinica*: 55(2): 174-181. (in Chinese with English abstract).
- [20] Zhu Dagang, Meng Xiangang, Zhao Xitao, et al, 2004. Evolution and climatic change of Nam Co of Tibet and ancient large lake in the northern Tibetan Plateau since the late Pleistocene [J]. *Geology in China*, 31(3): 269-277(in Chinese with English abstract).
- [21] Yafeng Shi, Ge Yu, Xiaodong Liu, Bingyuan Li and Tandong Yao, 2001. Reconstruction of the 30-40 ka BP enhanced Indian monsoon climate based on geological records from the Tibetan Plateau[J]. *Palaeogeography, Palaeoclimatology, Palaeoecology*: 169(1) pages 69-83.
- [22] Shi Yafeng, Jia Yulian, Ge Yu, et al, 2002. Features, impacts and causes of the high temperature and large precipitation event in the Tibetan Plateau and its adjacent area during 40-30 kaBP [J]. *Journal of Lake Science*: 14(1) pages 1-11 (in Chinese with English abstract).
- [23] Ge Yu, Lai Geyang, Jian Liu et al, 2003. Late MIS 3 climate simulations[J]. *Quaternary Sciences*: 23(1), 12-24 (in Chinese with English abstract)